Spatial correlation in ground motion intensities:

Measurement, prediction, and implications

Jack W. Baker

Thanks to former students Lukas Bodenmann, Yilin Chen, Nirmal Jayaram, Christophe Loth, Maryia Markhvida, Mahalia Miller

Thanks to collaborators Brendon Bradley, Božidar Stojadinović, Eric Thompson, David Wald, Bruce Warden

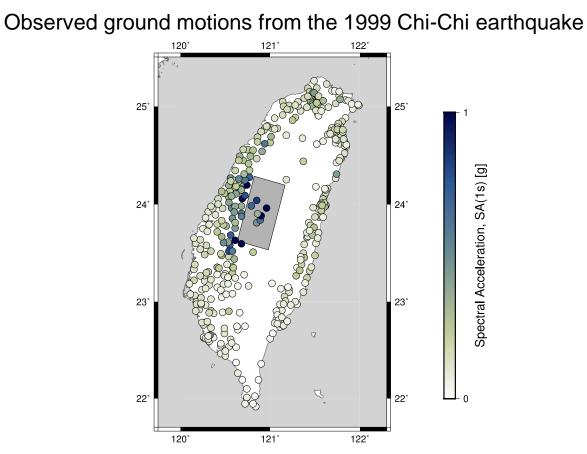


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Thank you Dimitrios!

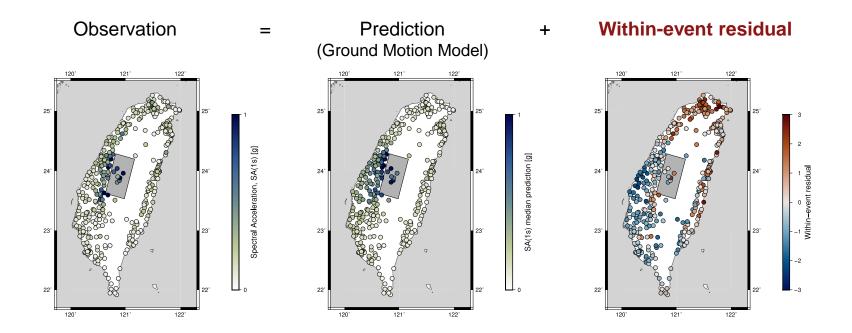


Spatial correlations in ground motion intensity



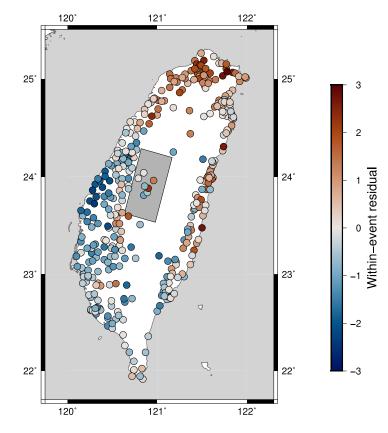
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Spatial correlation in ground motion

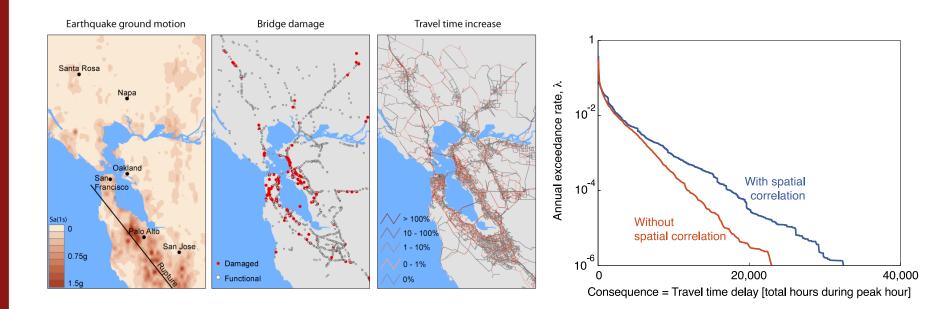


Correlation results from several factors

- Similar wave propagation paths
- Similar local site effects
- Similar location to rupture asperities

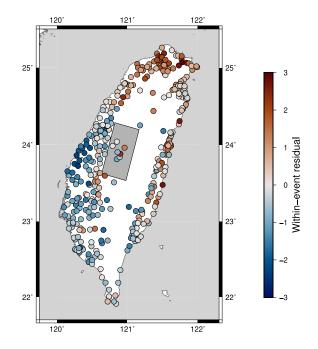


These correlations matter for regional risk simulation



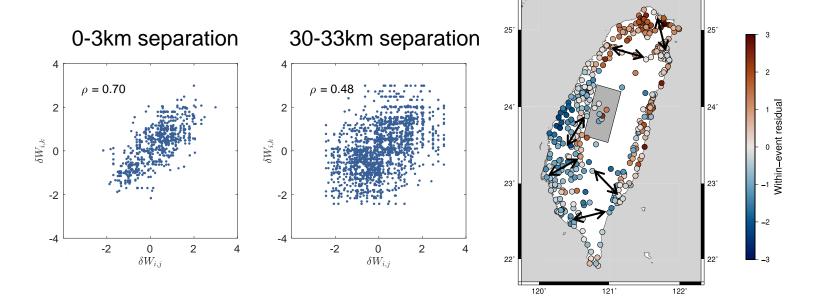
Estimation of spatial correlations

Ideally, we would like many observations for each station pair

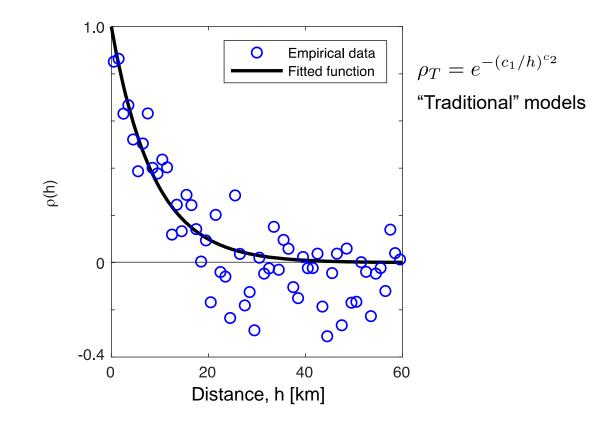


Estimation of spatial correlations

Typical practice: assume any pair of sites with equal separation distance has the same correlation ("stationarity")



Estimation of spatial correlations



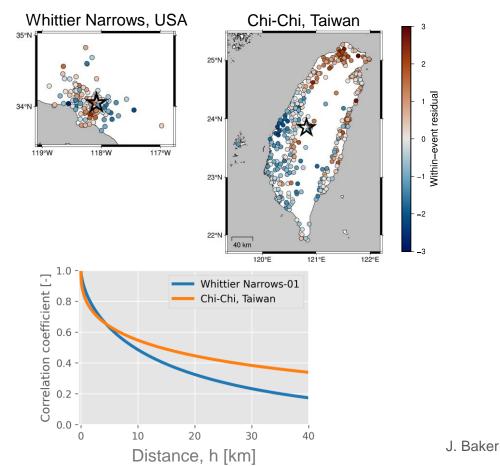
That's the basic story, but it misses some complexities...

Event-specific correlations, or something else?

Many studies have reported event-to-event variations

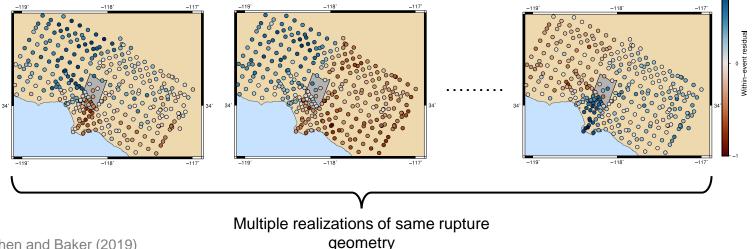
Hypothesized explanations:

- Regional variations
- Magnitude dependence
- Site conditions
- Event-specific variability



One path forward: study "physics-based" simulations

- We used CyberShake simulations (Graves et al. 2011) based on wave propagation through a 3D crustal velocity model
- Earthquake ruptures are described kinematically by slip amplitude, direction, and timing across the fault, and multiple realizations are available
- We find evidence of higher correlations in station pairs sharing a wave propagation path or similar geological conditions

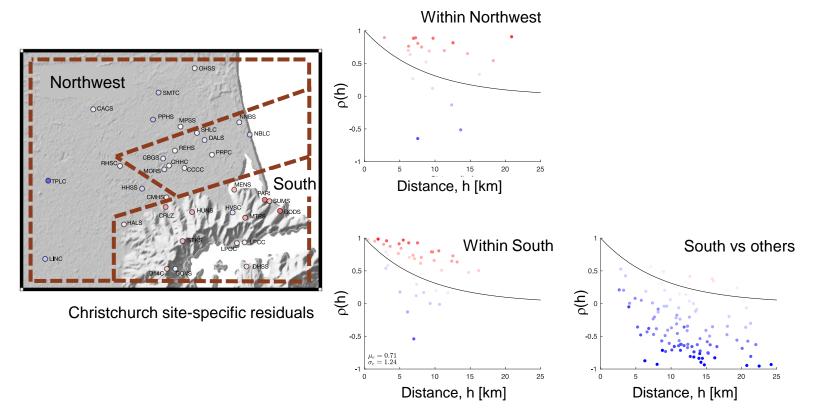


A second path forward: Stations with repeated observations

26 ground motion instruments 25 notable earthquakes -43°00' O OHSS O SMTC OCACS Christchurch Region NBLC -43°30' BHS OTPLC OLINC -44°00' 171°30' 172°00' 172°30' 173°00'

Christchurch, New Zealand

Correlations are stronger within geologic regions



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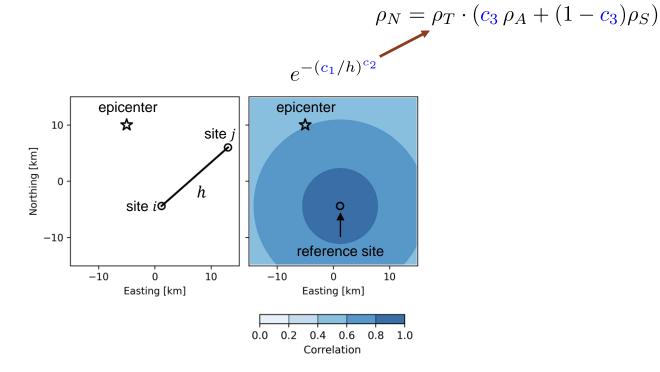
A new predictive model: Consider extra geometry and soil condition variability

Traditional models:

$$\rho_{\rm T} = e^{-(c_1/h)^{c_2}}$$

New model:

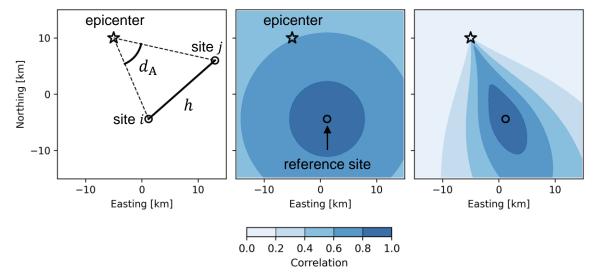
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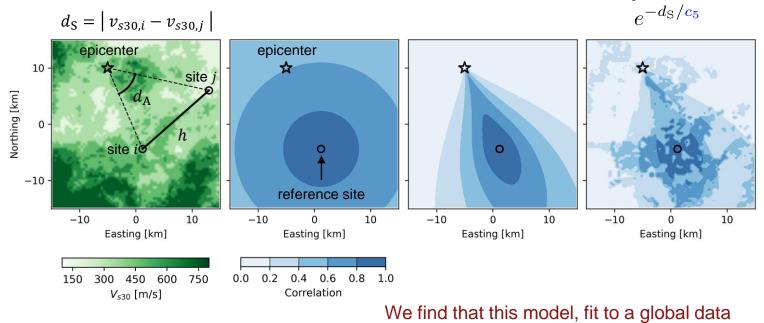
A new predictive model: Consider extra geometry and soil condition variability

$$\rho_N = \rho_T \cdot (c_3 \, \rho_A + (1 - c_3) \rho_S)$$

$$(1 + d_A/c_4)(1 - d_A/180)^{180/c_4}$$



A new predictive model: Consider extra geometry and soil condition variability $\rho_N = \rho_T \cdot (c_3 \rho_A + (1 - c_3)\rho_S)$



Bodenmann, Baker, Stojadinović (2023)

We find that this model, fit to a global data set, is more predictive than fitting traditional models to individual J. Baker

Predictive power versus model complexity

Traditional models (2003-):

 $\rho_{\rm T} = e^{-(c_1/h)^{c_2}}$

Compare to ground motion models Esteva and Rosenbluth (1964)

 $SA = c_1 + c_2M - c_3\ln R$

Predictive power versus model complexity

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New model:

 Boore, Joyner, Fumal (1997)

$$\mu_{\ln SA} = c_0 + c_1 (M - 6) + c_2 (M - 6)^2 + c_3 \ln \left(\sqrt{R^2 + c_4^2}\right) + c_5 \ln(V_{S,30})$$

Predictive power versus model complexity

Traditional models:

$$\rho_{\rm T} = e^{-(c_1/h)^{c_2}}$$

New model: $\rho_N = \rho_T \cdot (c_3 \rho_A + (1 - c_3) \rho_S)$ $e^{-(c_{1}/h)^{c_{2}}} e^{-d_{S}/c_{5}} Chiou and Youngs (2014)} e^{-(c_{1}/h)^{c_{2}}} e^{-d_{S}/c_{5}} (1+d_{A}/c_{4})(1-d_{A}/180)^{180/c_{4}}} + \frac{c_{u}}{c_{u}+c_{ooh(2-\max(M_{1}-4.5,0))}} e^{-c_{1}+\left\{c_{u}+\frac{c_{u}}{cosh(2-\max(M_{1}-4.5,0))}\right\}r_{sou}} + \frac{\ln(y_{o}) = \ln(y_{of_{1}}) + h}{h_{1}+c_{u}+\frac{c_{u}}{cosh(2-\max(M_{1}-4.5,0))}} e^{-c_{1}+\frac{c_{u}}{cosh(2-\max(M_{1}-4.5,0))}} e^{-c_{1}+\frac{c_{u}}{cosh(2-\max(M_{1}-4.5,0)}} e^{-c_{1}+\frac{c_{u}}{cosh(2-\max(M_{1}-4.5,0))$

Compare to ground motion models Esteva and Rosenbluth (1964)

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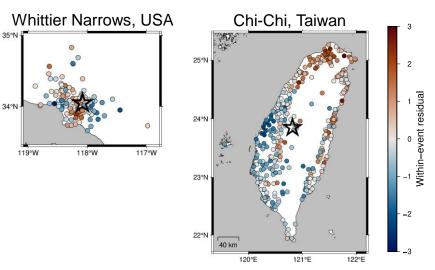
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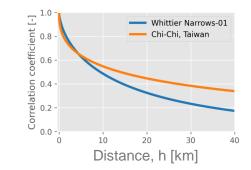
$$\begin{split} &\ln(y_{eq0}) = c_1 + \left\{ c_{1s} + \frac{c_{1s}}{\cosh(2 \cdot \max(\mathbf{M}_1 - 4.5, 0))} \right\} F_{gy_1} & \ln(y_0) = \ln(y_{eq_0}) + \eta_1 \\ &+ \left\{ c_{1s} + \frac{c_{2s}}{\cosh(2 \cdot \max(\mathbf{M}_1 - 4.5, 0))} \right\} F_{gy_0} & + \phi_1 \cdot \min\left(\ln\left(\frac{V_{Sy_0}}{1130}\right), 0 \right) \\ &+ \left\{ c_1 + \frac{c_{2s}}{\cosh(2 \cdot \max(\mathbf{M}_1 - 4.5, 0))} \right\} \Delta Z_{TOB} & + \phi_2(e^{\phi_1(\min(y_{2s}, 1130) - 360)} - e^{\phi_1(1130 - 360)}) \ln\left(\frac{y_{eq_1}e^{\phi_1} + \phi_4}{\phi_4} + \left\{ c_{11} + \frac{c_{11}}{\cosh(2 \cdot \max(\mathbf{M}_1 - 4.5, 0))} \right\} \Delta Z_{TOB} & + \phi_2(e^{\phi_1(\min(y_{2s}, 1130) - 360)} - e^{\phi_1(1130 - 360)}) \ln\left(\frac{y_{eq_1}e^{\phi_1} + \phi_4}{\phi_4} + \left\{ c_{11} + \frac{c_{11}}{\cosh(2 \cdot \cos(1 - e^{-\Delta Z_{10}}/\phi_1)} + e_g + e$$

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Event-specific correlations, or something else?



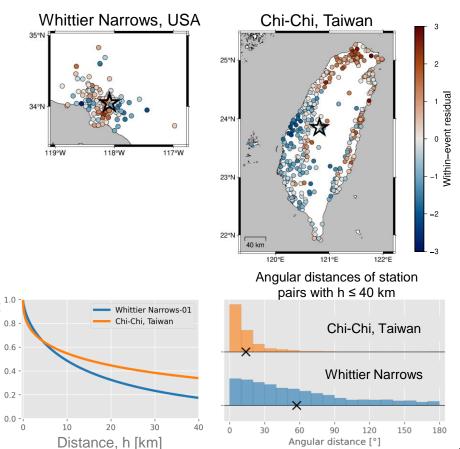
Many studies have reported event-to-event variations



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Event-specific correlations, or something else?

Correlation coefficient [-] 3.0 3.0 3.0 3.0 1



Many studies have reported event-to-event variations

Apparent explanations:

- Regional variations
- Magnitude dependence
- Source-to-site geometry
- Site conditions
- Event-specific variability

Conclusions

Ground motions exhibit spatially correlated amplitudes, and these correlations have important practical impacts

Well-recorded earthquakes and ground motion simulations give us a means to measure these correlations and build predictive models

The next generation of correlation models will:

- Utilize numerical simulations plus observations
- Account for the effects of path, site, ?

www.jackwbaker.com/joyner.htm