

DrinC: a software for drought analysis based on drought indices

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Abstract: Drought is a complex phenomenon which can be characterised mainly by its severity, duration and areal extent. Among these three dimensions, drought severity is the key factor which can be used for drought analysis. Drought indices are typically used to assess drought severity in a meaningful way. DrinC (Drought Indices Calculator) is a software package which was developed for providing a simple, though adaptable interface for the calculation of drought indices. The paper aims at presenting the overall design and the implementation of the software along with the utilisation of various approaches for drought analysis. DrinC can be used for the calculation of two recently developed indices, the Reconnaissance Drought Index (RDI) and the Streamflow Drought Index (SDI), as well as two widely known indices, the Standardised Precipitation Index (SPI) and the Precipitation Deciles (PD). Moreover, the software includes a module for the estimation of potential evapotranspiration (PET) through temperature based methods, useful for the calculation of RDI. The software may be used in a variety of applications, such as drought monitoring, assessment of the spatial distribution of drought, investigation of climatic and drought scenarios, etc. The applications of DrinC in several locations, especially in arid and semi-arid regions, show that it is gaining ground as a useful research and operational tool for drought analysis.

Key words: Drought analysis software, Drought indices, Reconnaissance drought index (RDI), Streamflow drought index (SDI), Standardized precipitation index (SPI), Potential evapotranspiration (PET)

1. INTRODUCTION

Drought is considered as a major natural hazard, affecting several sectors of the economy and the environment worldwide. It affects almost all the determinants of the hydrological cycle starting from precipitation and ending with streamflow in the surface water systems or the recharge and storage in the groundwater aquifers. Therefore, the determinant to be chosen depends upon which part of the hydrological cycle we are interested to focus our analysis (Tsakiris et al. 2013).

Conventionally, droughts are characterised by the determinant studied and are categorised as meteorological (precipitation, potential evapotranspiration), hydrological (streamflow or groundwater recharge), agricultural (soil moisture). Drought indices are essential tools for the characterisation and the monitoring of drought, since they simplify the complex climatic functions and can quantify climatic anomalies as for their severity, duration and frequency. They are also very useful as they can communicate to the wider audience easily comprehensible information regarding the severity of drought episodes (Tsakiris et al. 2007c). Drought indices can be used in the academic field (education, research, studies etc.), but also at strategic and operational levels (decision making, proactive management and drought mitigation etc.), having a wide range of potential users. Of course, no single index is universally suitable, so in most cases it is necessary to consider more than one index (Morid et al. 2006).

Drought indices are usually calculated, either by applying manually the corresponding equations and procedures, or by using tools designed for this purpose. Such tools may work within the framework of other software packages (e.g. MATLAB, R-project, ArcGIS, etc.), or as stand-alone software, including one or more indices. Indicatively, some of the available tools for drought indices calculation are:

- The SPI_SL_6 (NDMC 2014), a simple application for calculating the Standardized Precipitation Index for various monthly intervals.
- The SPATSIM, mainly a water resources information management software package, which can also be used for the calculation of precipitation based indices, such as the Standardized

Precipitation Index (SPI), the Deciles, the Effective Drought Index (EDI) and the departures from the long-term mean and median precipitation. It contains an integrated database management system that uses GIS shape files as the main form of data access values (Smakhtin and Hughes 2007).

- The SPEI package (Beguería and Vicente-Serrano 2013), which is a set of functions for R-project that can be used for computing the Standardized Precipitation-Evapotranspiration Index (SPEI) and the SPI.
- The CDI, a MS Excel-based software, developed by FAO SWALIM (Balint et al. 2011), which computes four indices: the Precipitation Drought Index (PDI), the Temperature Drought Index (TDI), the Vegetation Drought Index (VDI) and the Combined Drought Index (CDI).

Many of the existing tools often require advanced computer skills by the user or familiarity with the specific software packages, restricting the task of calculation of drought indices to experts. The development of specialised, stand-alone software for the calculation of drought indices, that provides a simple interface and produces clear and comparable outputs, can be very helpful for both academic and operational use. Further, the consideration of the results of several drought indices may lead to a comprehensive assessment of drought severity and give a better perspective to specific research or management objectives.

On the basis of the above considerations, the software package named DrinC was developed at the Centre for the Assessment of Natural Hazards and Proactive Planning and the Laboratory of Reclamation Works and Water Resources Management of the National Technical University of Athens. It can be used for the calculation of drought indices - and especially of two that were recently developed - suitable for drought characterisation, drought monitoring, spatial analysis of drought and investigation of climatic scenarios. The user can select, through a windows-based interface, among several options that may fit better to the specific requirements of each task or case study. Further, DrinC includes two modules for the calculation of potential evapotranspiration through temperature-based methods and for drought monitoring, respectively.

In this paper, the objectives, the overall design and the implementation of DrinC software are presented. Moreover, various approaches for drought analysis and their utilisation by the software are described, along with some real-world applications.

2. DROUGHT INDICES OVERVIEW

DrinC (Drought Indices Calculator) aims at providing a user-friendly tool for the calculation of several drought indices. Key objective in its design was the widest possible applicability for several types of drought (meteorological, hydrological, agricultural) and different locations. It was also taken into account that drought studies are particularly essential in arid and semi-arid regions, where data availability is usually limited. Therefore, the main criteria for the selection of indices were: a) to have relatively low data requirements, allowing the application of the software in many regions and b) their results to be clearly interpreted for direct and efficient operational use.

Based on these criteria, two recently developed and two more widely known indices were included in DrinC: the Reconnaissance Drought Index (RDI), the Streamflow Drought Index (SDI), the Standardised Precipitation Index (SPI) and the Precipitation Deciles (PD).

As can be easily understood, RDI, SPI and PD refer to the meteorological drought and use as main determinant the precipitation (and additionally the potential evapotranspiration for RDI only). Further, RDI can also be used for the agricultural drought analysis, as it can adequately describe the water balance, and it is particularly useful when reference periods related to development stages of the crop are selected (Tsakiris et al. 2010). On the other hand, SDI applies to hydrological drought and uses the streamflow as the key determinant.

Apart from the originally proposed methods of calculation for each index, DrinC incorporates alternative methods that allow the comparison of the results among the indices. Further, this gives a

key advantage to the user, since it provides the flexibility to select among various options for adjusting the outputs to his particular needs.

Following, there is a short presentation and the key characteristics of the drought indices calculated by DrinC.

2.1 Precipitation Deciles

One of the simplest meteorological drought indices is the method of Deciles which was introduced by Gibbs and Maher (1967). The precipitation totals for the preceding three months are ranked against climatologic records and if the sum falls within the lowest decile of the historical distribution of 3-month totals, then the region is considered to be under drought conditions (Kininmonth et al. 2000). The drought ends when: (i) the precipitation measured during the past month already places the 3-month total in or above the fourth decile, or (ii) the precipitation total for the past three months is in or above the eighth decile. The advantage of the method of deciles is its computational ease, but its simplicity can lead to conceptual difficulties. The deciles are grouped into five classes as presented in Table 1.

Table 1. Classification of drought conditions according to deciles.

Decile Class	Description
deciles 1-2: lowest 20%	much below normal
deciles 3-4: next lowest 20%	below normal
deciles 5-6: middle 20%	near normal
deciles 7-8: next highest 20%	above normal
deciles 9-10: highest 20%	much above normal

2.2 Reconnaissance Drought Index (RDI)

The Reconnaissance Drought Index (RDI) was developed to approach the water deficit in a more accurate way, as a sort of balance between input and output in a water system (Tsakiris and Vangelis 2005; Tsakiris et al. 2007c). It is based both on cumulative precipitation (P) and potential evapotranspiration (PET), which are one measured (P) and one calculated (PET) determinant. The initial value (α_k) of RDI is calculated for the i -th year in a time basis of k (months) as follows:

$$\alpha_k^{(i)} = \frac{\sum_{j=1}^k P_{ij}}{\sum_{j=1}^k PET_{ij}}, \quad i = 1(1)N \quad \text{and} \quad j = 1(1)k \quad (1)$$

in which P_{ij} and PET_{ij} are the precipitation and potential evapotranspiration of the j -th month of the i -th year and N is the total number of years of the available data.

The values of α_k follow satisfactorily both the lognormal and the gamma distributions in a wide range of locations and different time scales, in which they were tested (Tigkas 2008; Tsakiris et al. 2008). By assuming that the lognormal distribution is applied, the following equation can be used for the calculation of RDI_{st} :

$$RDI_{st}^{(i)} = \frac{y^{(i)} - \bar{y}}{\hat{\sigma}_y} \quad (2)$$

in which $y^{(i)}$ is the $\ln(\alpha_k^{(i)})$, \bar{y} is its arithmetic mean and $\hat{\sigma}_y$ is its standard deviation.

In case the gamma distribution is applied, the RDI_{st} can be calculated by fitting the gamma

probability density function (pdf) to the given frequency distribution of α_k (Tsakiris et al. 2008; Tigkas 2008). For short reference periods (e.g. monthly or 3-months) which may include zero values for the cumulative precipitation of the period, the RDI_{st} can be calculated based on a composite cumulative distribution function including the probability of zero precipitation and the gamma cumulative probability.

Positive values of RDI_{st} indicate wet periods, while negative values indicate dry periods compared with the normal conditions of the area. Drought severity can be categorised in mild, moderate, severe and extreme classes, with corresponding boundary values of RDI_{st} (-0.5 to -1.0), (-1.0 to -1.5), (-1.5 to -2.0) and (< -2.0), respectively.

2.3 Streamflow Drought Index (SDI)

According to Nalbantis (2008), if a time series of monthly streamflow volumes $Q_{i,j}$ is available, in which i denotes the hydrological year and j the month within that hydrological year ($j = 1$ for October and $j = 12$ for September), $V_{i,k}$ can be obtained based on the equation:

$$V_{i,k} = \sum_{j=1}^{3k} Q_{i,j} \quad i = 1, 2, \dots \quad j = 1, 2, \dots, 12 \quad k = 1, 2, 3, 4 \quad (3)$$

in which $V_{i,k}$ is the cumulative streamflow volume for the i -th hydrological year and the k -th reference period, $k = 1$ for October-December, $k = 2$ for October-March, $k = 3$ for October-June, and $k = 4$ for October-September.

Based on the cumulative streamflow volumes $V_{i,k}$, the Streamflow Drought Index (SDI) is defined for each reference period k of the i -th hydrological year as follows:

$$SDI_{i,k} = \frac{V_{i,k} - \bar{V}_k}{s_k} \quad i = 1, 2, \dots, \quad k = 1, 2, 3, 4 \quad (4)$$

in which \bar{V}_k and s_k are respectively the mean and the standard deviation of cumulative streamflow volumes of the reference period k as these are estimated over a long period of time. In this definition the truncation level is set to \bar{V}_k although other values based on rational criteria could be also used.

According to Nalbantis and Tsakiris (2009), states (classes) of hydrological drought are defined for SDI in an identical way to those used in the meteorological drought indices SPI and RDI. Five states are considered, which are denoted by an integer number ranging from 0 (non-drought) to 4 (extreme drought) and are defined through the criteria of Table 2.

Table 2. Definition of states of hydrological drought with the aid of SDI.

State	Description	Criterion
0	Non-drought	$SDI \geq 0.0$
1	Mild drought	$-1.0 \leq SDI < 0.0$
2	Moderate drought	$-1.5 \leq SDI < -1.0$
3	Severe drought	$-2.0 \leq SDI < -1.5$
4	Extreme drought	$SDI < -2.0$

2.4 Standardized Precipitation Index (SPI)

For the SPI calculation, the long-term precipitation record for a desired period is fitted to a probability distribution, which is then transformed into a normal distribution so that the mean SPI for the location and desired period is zero (McKee et al. 1993; Edwards and McKee 1997). Positive

SPI values indicate greater than median precipitation, and negative values indicate less than median precipitation. Since SPI is normalised, wetter and drier climates can be represented in the same way.

Thom (1958) found the gamma distribution to fit well to the climatological precipitation time series. The gamma distribution is defined by its frequency or probability density function:

$$g(x) = \frac{1}{\beta^a \Gamma(a)} x^{a-1} e^{-x/\beta}, \quad \text{for } x > 0 \quad (7)$$

in which α and β are the shape and scale parameters respectively, x is the precipitation amount and $\Gamma(\alpha)$ is the gamma function. Parameters α and β of the gamma pdf are estimated for each station and for each time scale of interest (1, 3, 6, 9, 12 months, etc.). Maximum likelihood estimations of α and β are:

$$a = \frac{1}{4A} \left(1 + \sqrt{1 + \frac{4A}{3}} \right), \quad \beta = \frac{\bar{x}}{a}, \quad \text{where } A = \ln(\bar{x}) - \frac{\sum \ln(x)}{n} \quad (8)$$

and n is the number of observations.

The resulting parameters are then used to find the cumulative probability of an observed precipitation event for the given month and time scale for the location in question. Since the gamma function is undefined for $x = 0$ and a precipitation distribution may contain zeros, the cumulative probability becomes:

$$H(x) = q + (1 - q)G(x) \quad (9)$$

in which q is the probability of zero precipitation and $G(x)$ is the cumulative probability of the incomplete gamma function. If m is the number of zeros in a precipitation time series, then q can be estimated by m/n . The cumulative probability $H(x)$, is then transformed to the standard normal random variable z with mean zero and variance of one (Abramowitz and Stegun 1965), which is the value of the SPI.

According to the SPI, a drought event occurs when the index continuously reaches an intensity of -1.0 or less. The event ends when the SPI becomes positive. Each drought event, therefore, has a duration defined by its beginning and end, and intensity for each month that the event continues. Drought magnitude is the positive sum of the SPI for each month during the drought event (Hayes et al. 2007).

Generally, monthly precipitation is not normally distributed so a transformation is performed such that the derived SPI values follow a normal distribution. The SPI is the number of standard deviations that the observed value would deviate from the long-term mean, for a normally distributed random variable. One interpretation of the resultant values is presented in Table 3 (Tsakiris and Vangelis 2004).

Table 3. Classification of drought conditions according to the SPI.

SPI values	Classification
2.0 or more	Extremely Wet
1.5 to 1.99	Very Wet
1.0 to 1.49	Moderately Wet
-.99 to .99	Near Normal
-1.0 to -1.49	Moderately Dry
-1.5 to -1.99	Severely Dry
-2 or less	Extremely Dry

3. SOFTWARE DESIGN AND IMPLEMENTATION

3.1 Software structure, interface and data management

DrinC was programmed in Visual Basic 6. During the development of DrinC, emphasis has been given on maintaining a simple, comprehensive and user-friendly structure. The flow chart of the software appears in Fig. 1.

DrinC has full graphical user interface functionality (GUI) and runs on MS Windows operating systems (Fig. 2). All the main functions can be accessed through the main menu of the software, while some additional options are available for specific operations.

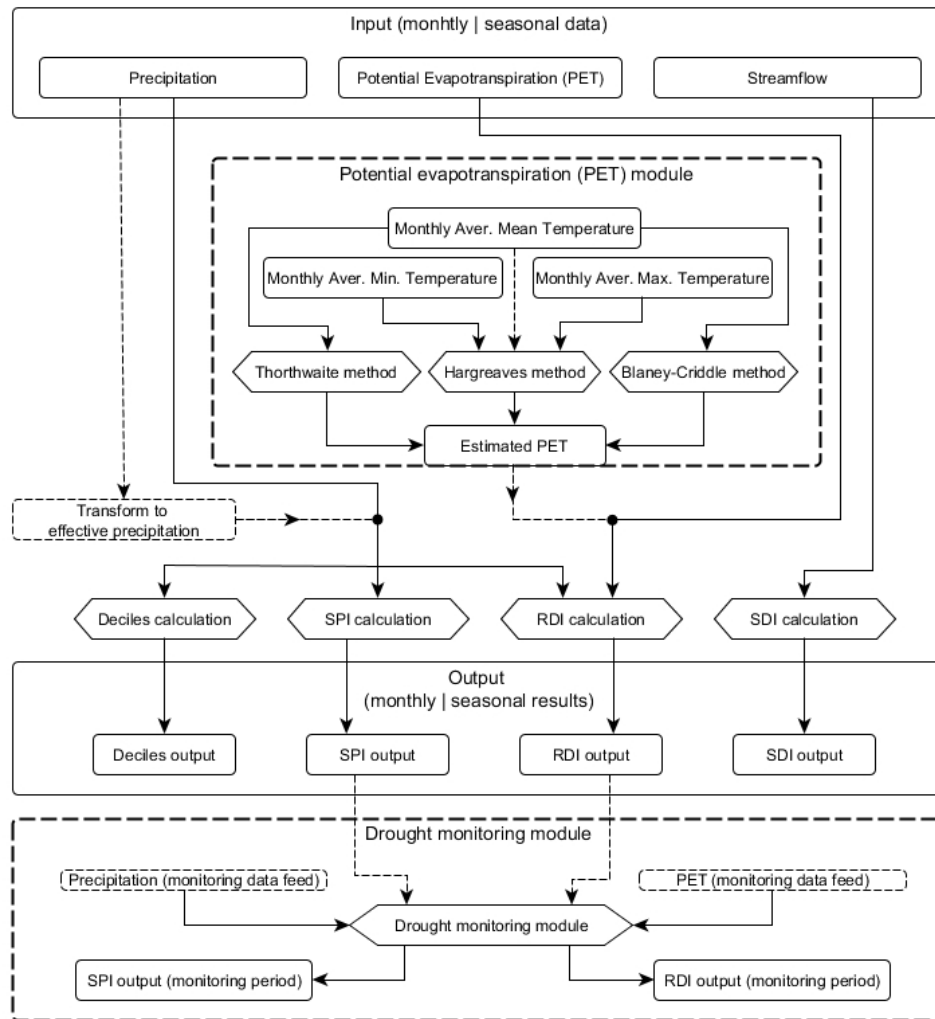


Fig. 1. The flow chart of DrinC software.

The calculation of each index requires different input data, as can be seen in Table 4. Data can be monthly, annual or seasonal and there is an option to use real (up to 150 years) or synthetic (up to 1500 years) data series. Data can be imported directly from MS Excel files (.xls). For monthly data the software automatically inspects the structure of the file and identifies the position of the data within the spreadsheet, while for annual or seasonal data the user defines the cells from which the loading should start. Also, there is the option to transform precipitation data to effective precipitation data, using either the U.S. Bureau of Reclamation method (Stamm 1967) or the F.A.O. method (Brouwer and Heibloem 1986). The imported data are loaded in a grid form where they can be reviewed.

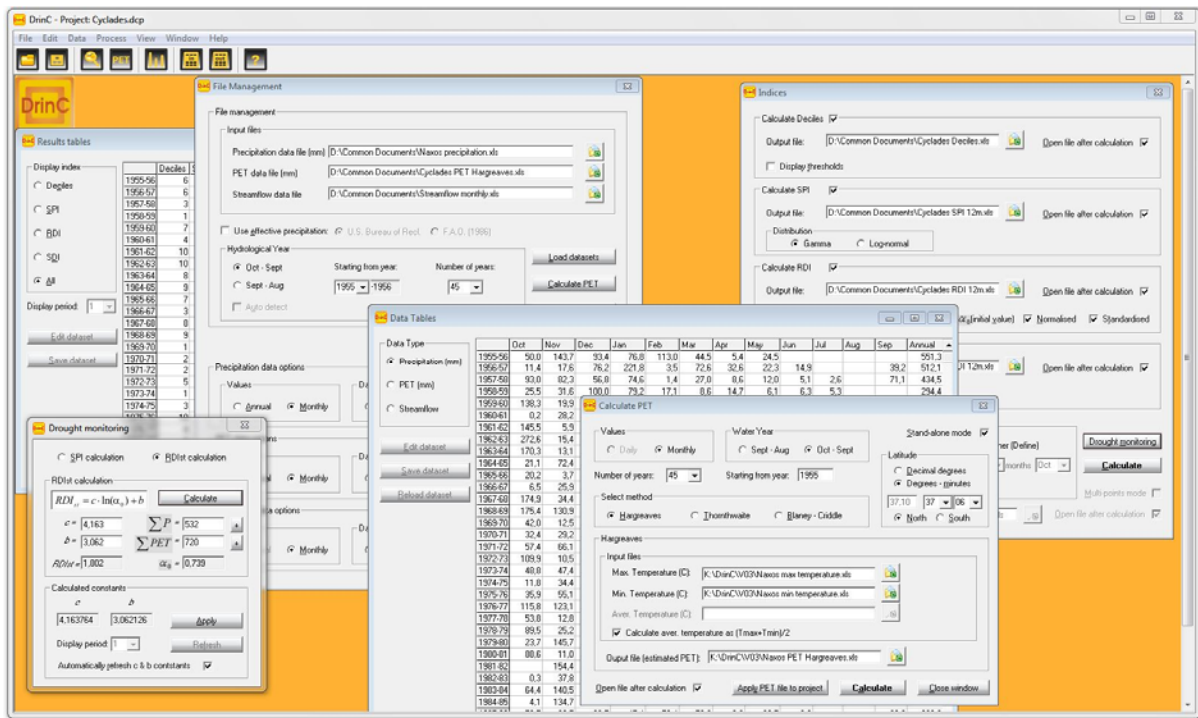


Fig. 2. The graphical user interface (GUI) of DrinC.

Table 4. Required input data for each index.

Index	Required input data
Deciles	Precipitation
RDI	Precipitation, PET (or temperature)
SPI	Precipitation
SDI	Streamflow

3.2 Drought indices calculation

The initial theoretical concepts of the SPI and the RDI are different, regarding their calculation procedure. The SPI is calculated as a “rolling” index in a monthly sequential order (McKee et al. 1993), usually with five running time intervals (i.e. 1-, 3-, 6-, 9- and 12-months). The RDI is computed for discrete reference periods (e.g. 3, 6, 9 and 12 months) on hydrological year basis (Tsakiris et al. 2007c). If these different approaches are used, the ability to compare the results of the indices is limited. Further, there is narrow flexibility to use them in several specific case studies, for example, if there is a need to focus on certain periods of the year.

In DrinC, there is the option to use both of the above approaches, giving the ability to formulate a drought analysis that suits better the needs or the purpose of a study or application. The results for each drought index may be presented per month or per period that allows a direct comparison of drought severity for specific periods of the year.

The primary reference base in DrinC is the hydrological year (October – September), therefore the default calculation period starts from October and the main calculation steps are 1-month, 3-months, 6-months and annual. However, it is possible to define other calculation steps (e.g. 4-months) or different starting months (e.g. January) through drop down lists and radio buttons. This flexibility may be useful for several real-world applications; for instance, the study of drought effects on specific crops should coincide to the crop growth period.

The results can be saved in MS Excel files (in one file or in separate files for each index) and the user may select to draw automatically the graph in the output file (Fig. 3).

Another useful feature is the ‘multi-points mode’ function, with which a set of annual or seasonal input data for several spatial points can be loaded and each drought index can be calculated

for each point through an iterative routine. This automated procedure for the assessment of spatial distribution of drought indices is particularly useful for the production of drought maps.

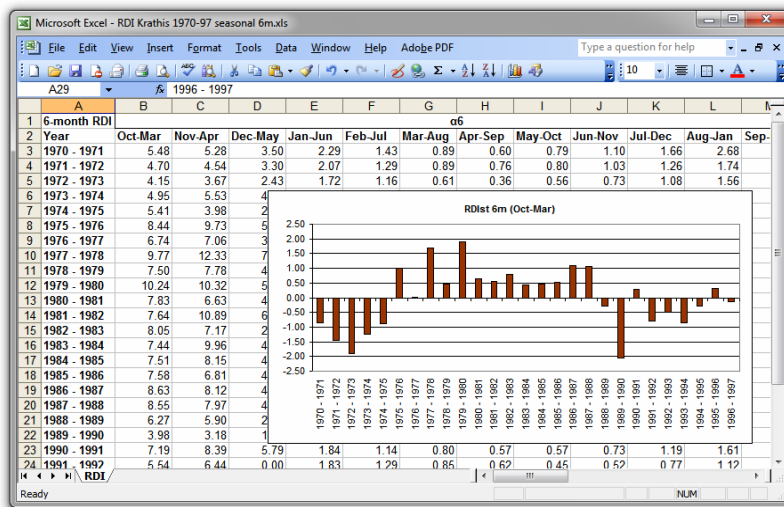


Figure 3. Example of an output file of DrinC.

3.3 Modules for PET calculation and drought monitoring

There are several methods to estimate PET, though some of those can be data demanding. However, according to Vangelis et al. (2013), temperature based methods to estimate PET would give sufficient results for the RDI calculation.

DrinC provides a module (Fig. 4) for the calculation of PET with three temperature based methods: Hargreaves (Hargreaves and Samani 1985), Thornthwaite (1948) and Blaney-Criddle (Doorenbos & Pruitt 1977). From these methods, the first requires minimum and maximum temperature data (mean temperature is optional), whereas the other two require only mean temperature data. The basic inputs that should be defined are the number of years, the starting year (leap years are taken into account) and the latitude.

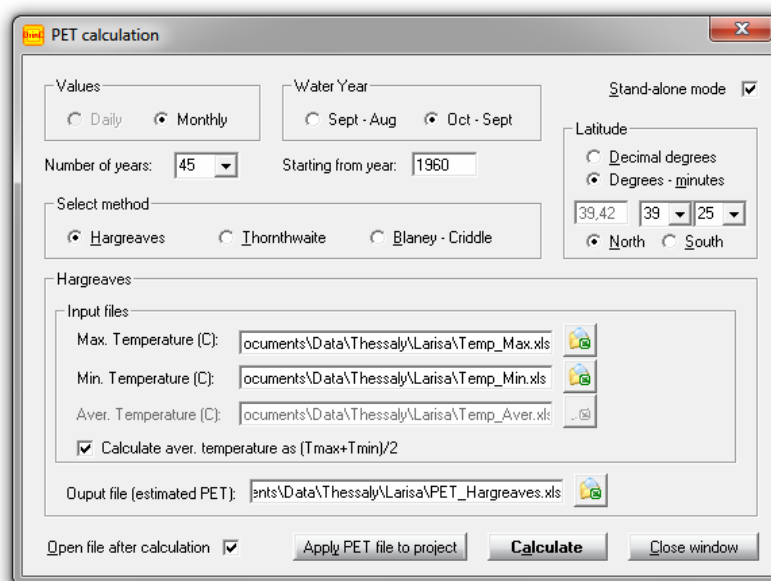


Fig. 4. The PET calculation module.

The drought monitoring module of DrinC can be used to directly estimate the drought indices for a present period, provided that there are sufficient available historical data in order to depict the climatic behaviour for the area of study. There is the assumption that the climatic characteristics of the study area remain the same for both the monitoring period and the available historical period.

The procedure is based on the linear function, which in case of RDI is:

$$RDI_{st} = c \cdot \ln(\alpha_n) + b \quad (10)$$

in which α_n is the initial value of RDI for n -months period, c and b are constants calculated from the available historical data for the corresponding n -months period.

4. DROUGHT ANALYSIS

4.1 Drought characterisation

Drought characterisation mainly focuses on the analysis of the historical drought conditions of an area. From a stochastic point of view, it provides information required for the subsequent risk analysis. Also, it provides spatial and temporal representations of historical droughts and therefore places current conditions in historical perspective. Drought characterisation can be valuable for the evaluation of the risk of future droughts, for the determination of the impacts of drought on various sectors (economy, environment, society etc.) and for the design of the preparedness plans for the management of the anticipated drought impacts (Rossi et al. 1992).

Drought indices are widely accepted as useful elements to characterise drought and its statistical properties. Some basic steps for drought characterisation can be described as follows (Tsakiris et al. 2007a):

- i. Collect all the necessary monthly data of each reliable meteorological station of the selected area
- ii. Transfer meteorological data at a basin or sub-basin level
- iii. Calculate selected indices on an appropriate time scale (e.g. trimester, semester or annual scale)
- iv. Display the results graphically using colours for the various levels of drought intensity
- v. Analyse the frequency of occurrence of 1-year drought using historical data. Other time steps could be also used.

Some of the above parameters (area, time scale etc.) should be carefully selected, based on the system under study and the purpose of the characterisation. For instance, an annual time scale is appropriate for assessing the impact of drought on a water body or a reservoir, while shorter time scales (e.g. 1-month or 3-months) could be used for the effects of drought on agricultural production.

DrinC provide all the above options for drought indices calculation. For example, the drought conditions of an area for 45 hydrological years are shown in Fig. 5. Each drought class is represented by a different colour for a better visualisation of drought intensity; e.g. the severe (1958-59 and 1992-93) and the extreme (1989-90 and 1999-2000) events are illustrated with orange and red colours, respectively. Also, Figs. 6 and 7 present a comparative analysis of the frequency of drought occurrences and the number of drought events along with their duration, respectively, for four areas of Greece (period 1955-2002).

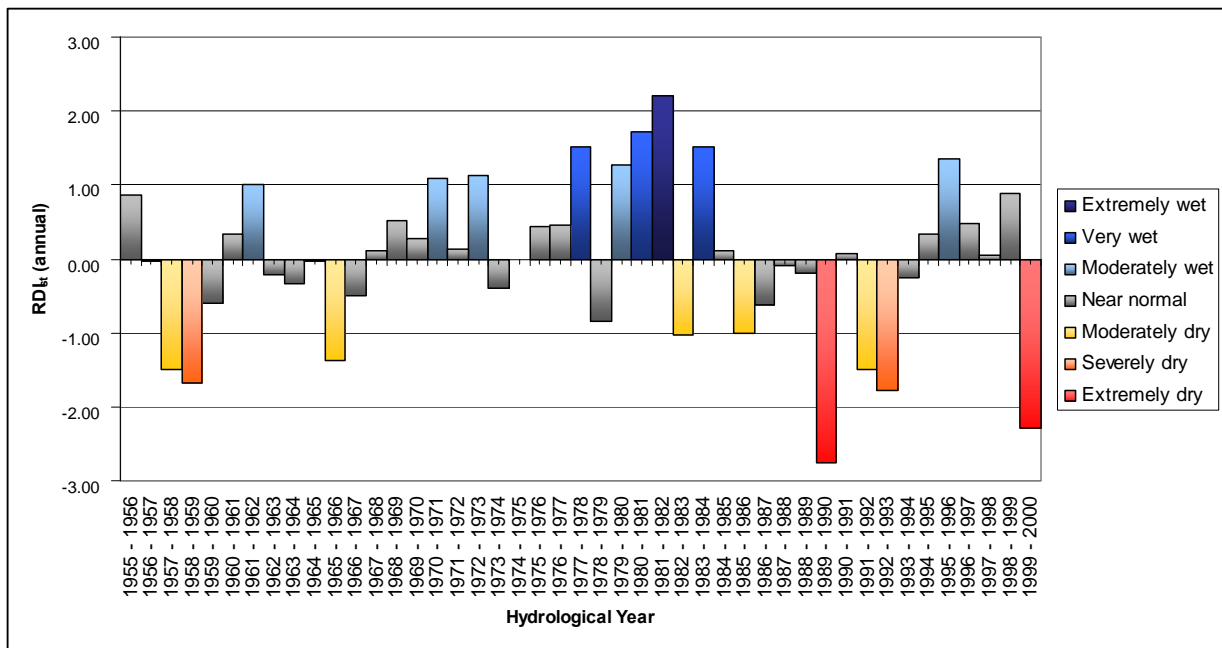


Fig. 5. Drought conditions of a region (Naxos, Greece) expressed by the annual RDI_{st} for 45 hydrological years (different colours represent each drought class).

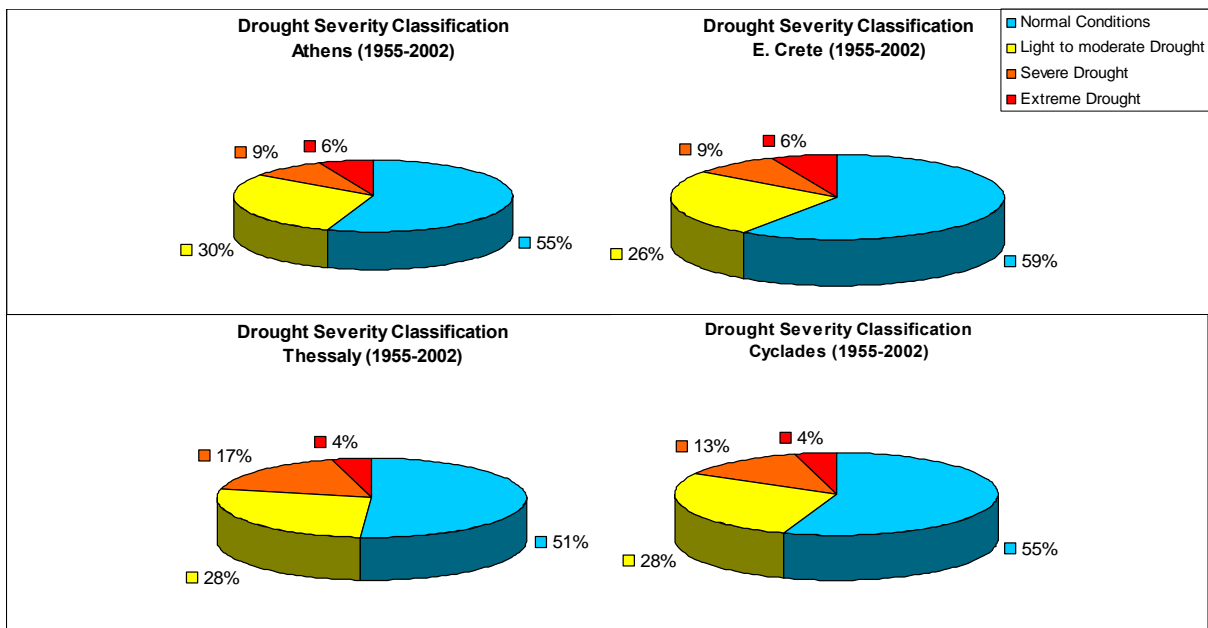


Fig. 6. Frequency of droughts for four areas of Greece (period 1955-2002) (Tigkas 2008).

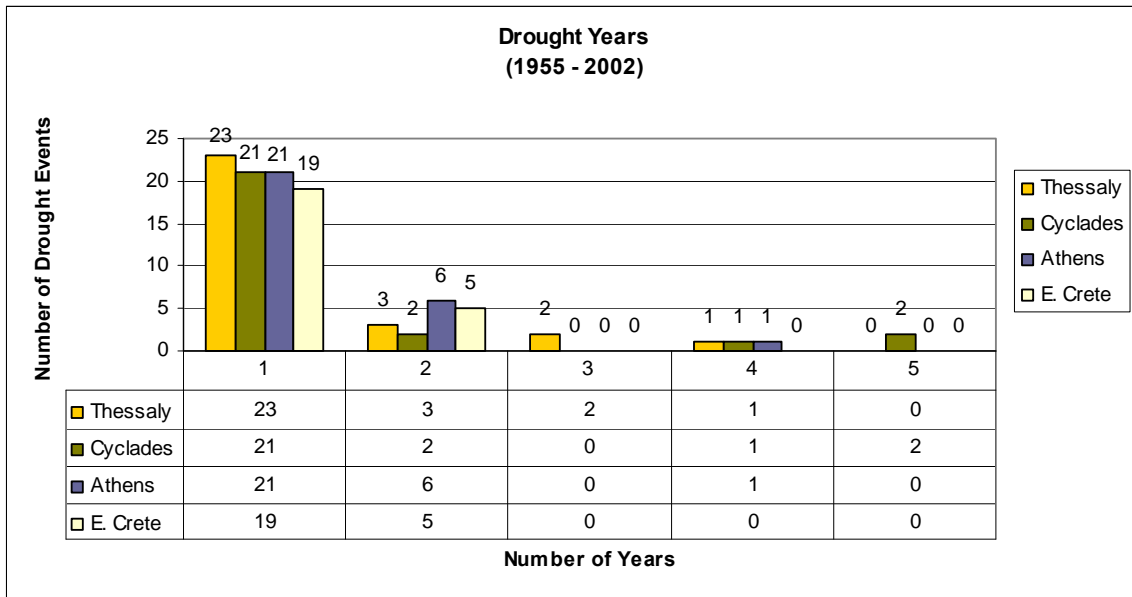


Fig. 7. Duration of consecutive drought years for four areas of Greece (period 1955-2002) (Tigkas 2008).

4.2 Drought Monitoring

Drought monitoring is essential for establishing an early warning system and for adopting the appropriate measures for the mitigation of the anticipated damages (Wilhite et al. 2005; Wilhite 2009; Bordi and Sutera 2007).

The main variable involved in drought analysis is precipitation, while potential evapotranspiration (PET) may also be used (e.g. in RDI). Therefore, the design of a meteorological drought monitoring network could be based on precipitation and air temperature (which is needed for PET estimation). However, since the spatial variability of precipitation generally exceeds that of temperature, the crucial variable for network design can be the precipitation only (Tsakiris et al. 2008). For establishing the drought monitoring system it is necessary to have available the historical data series that provides the climatic context of the area, while the new data should be regularly provided, usually on monthly basis (Tsakiris and Pangalou 2009).

DrinC can be used to monitor drought either for a predefined reference period (e.g. a “rolling” 3-month period) or for various reference periods starting from the same month (e.g. ‘Oct-Dec’, ‘Oct-Jan’, ‘Oct-Feb’, etc.), according to what fits better to the monitoring purpose.

4.3 Assessment of the spatial distribution of drought

Although there are a lot of studies which propose simplifications in drought analysis mainly regarding the drought dimension of areal extent (Tsakiris et al. 2007b), this drought dimension remains an important aspect in many drought studies. A significant simplification towards this direction is the use of the base hydrological area, which is the catchment area or river basin, instead of a discretised mesh covering the area of interest.

However, in many cases the areal extent of drought should be approached by a discretised model as the one presented in Fig. 8. This kind of models can be particularly useful, especially when remote sensing methods are employed for the calculation of drought indices for spatio-temporal drought analyses (e.g. Dalezios et al. 2012; Kanellou et al. 2012). For such cases, DrinC provides a feature which helps the user compute the required drought indices for the entire set of cells of the grid (mesh) simultaneously, avoiding the repetition of the procedure for thousands of times for each grid cell. The final results can be illustrated in thematic maps of the area, providing a clear picture of the different drought severity zones of the area.

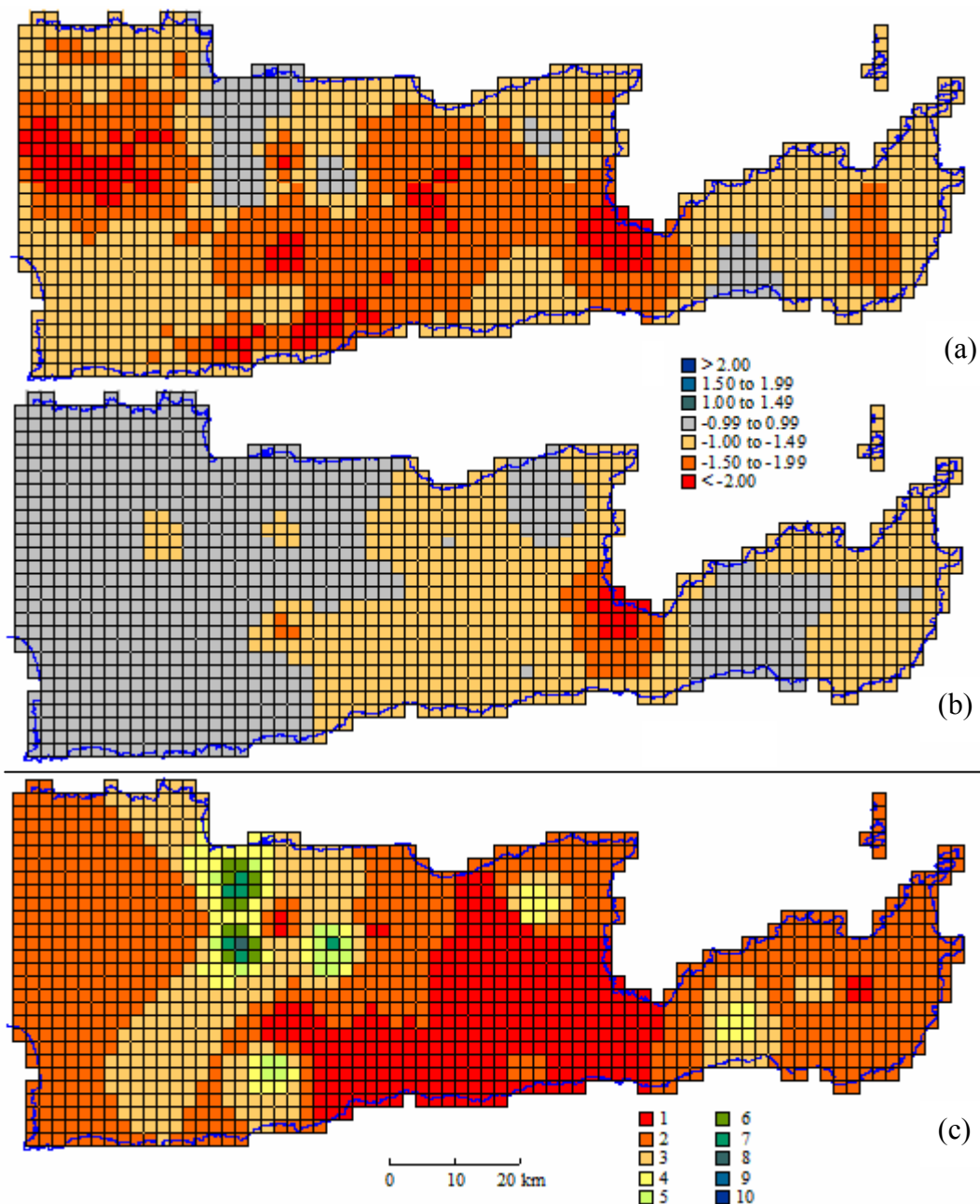


Fig. 8. Annual drought maps (Eastern Crete island – Greece) presenting the spatial distribution of drought with the use of a) RDI, b) SPI and c) Deciles.

4.4 Drought and climatic trends and scenarios

The RDI, apart from its use as a drought index, can also be used as a climatic index, suitable to identify possible climatic trends in annual or seasonal basis. For example, Fig. 9 presents the initial value of RDI (α_{12}) for two geographical areas in different climatic zones: the first for an area with humid climate, the second for an area with semi-arid climate. It can be observed that α_{12} is significantly higher for the humid area for the entire timeseries associated with higher variability in comparison with the semi-arid area. In order to detect the trend in the timeseries of annual RDI (α_{12}) one can use all appropriate methods for trend detection (e.g. Spearman test, Mann-Kendall test, Runs test, Durbin-Watson test, etc.). If the trend is statistically significant, then it can be assumed that there are strong indications of climatic change at annual basis (Tigkas et al. 2013).

Also, the RDI can serve for the investigation of the drought conditions for climate change scenarios. In this context, DrinC can produce series of results based on real or synthetic data in

order to evaluate the possible future climatic and drought conditions of a region. Further, this can be used for the assessment of the drought impacts on the economy, the society and the environment (e.g. crop production, drinking water availability, streamflow changes).

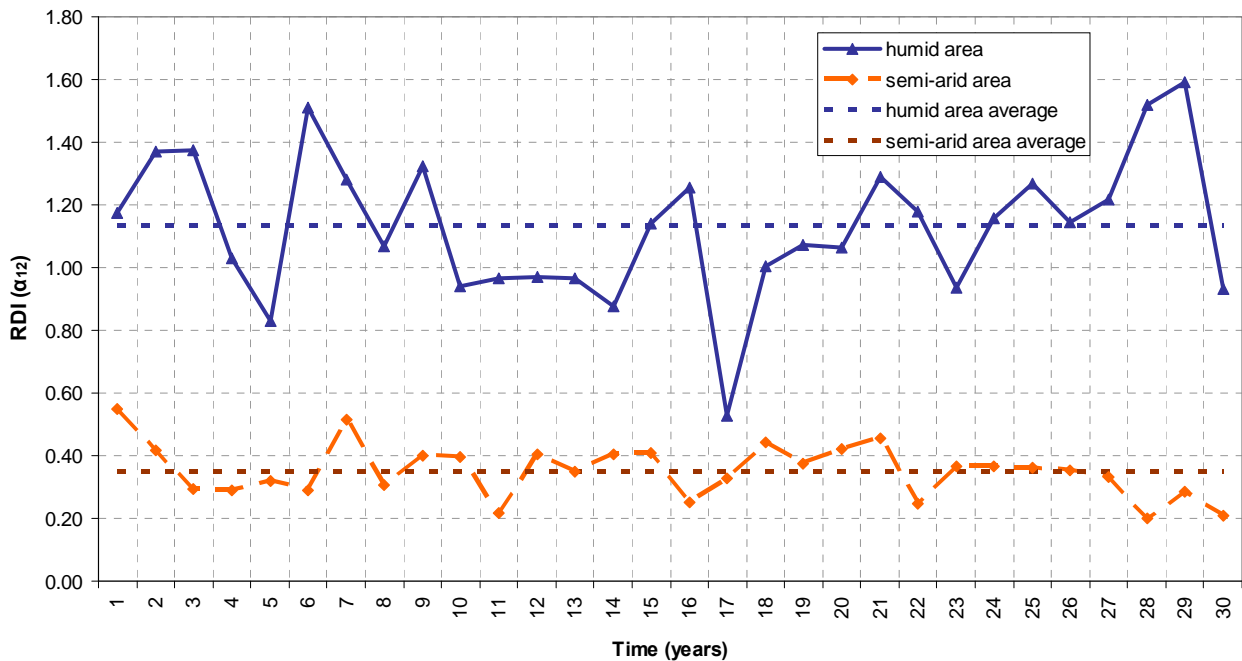


Fig. 9. The initial value (α_{12}) of RDI for a time series of 30 years for a humid (Aspropotamos, Greece) and a semi-arid area (Heraklion, Greece) (dashed line represents the average of α_{12} for each area).

4.5 Real-world applications

Several institutions and research groups have used DrinC in various applications and drought studies, mainly in arid and semi-arid regions. A few examples, that illustrate the range of applications that DrinC can be used, are presented below:

- Greece – Crete: The Centre for the Assessment of Natural Hazards and Proactive Planning of the National Technical University of Athens (NTUA) studied drought episodes on a regional level at the island of Crete. A series of annual drought maps were produced with the use of DrinC along with a spatial distribution model (Tsakiris et al. 2007d).
- Italy: The Water Observatory of the Region of Sicily and the University of Palermo analysed time series of climatic data (rainfall, temperature) using drought indices (SPI, RDI, NDVI) and vegetation data at the Oreto watershed in Sicily. Both satellite and conventional measurements were used in order to assess climatic effects on vegetation at the study area (Capodici et al. 2008).
- Cyprus: The Meteorological Service of Cyprus has used DrinC in two studies for the assessment and the monitoring of drought in Cyprus, using an extensive network of meteorological stations and GIS techniques (Pashiardis and Michaelides 2008; Michaelides and Pashiardis 2008).
- Greece – Thessaly: The University of Thessaly used DrinC to assess meteorological drought for a period of 45 years, using three drought indices (Deciles, Palmer Z and RDI) (Kanellou et al. 2008).
- Malta: The Water Directorate of the Malta Resources Authority used DrinC to analyse drought events in the country through RDI and SPI. Also, there was an attempt to assess the problems that could be caused by water shortage due to severe drought events (Borg 2009).
- Iran: The University of Isfahan and the Meteorological Organisation of Iran used DrinC in a study for the operational drought monitoring in the region of Isfahan. Meteorological data

- from 10 stations were used for the calculation of RDI and SPI (Mostafavi Darani et al. 2011).
- Greece – Peloponnese: The Lab. of Reclamation Works and Water Resources Management of the NTUA used DrinC along with Medbasin rainfall-runoff model (Tigkas and Tsakiris 2004) in order to assess the drought and climatic change impact on streamflow in small watersheds. The case study was performed in 6 small watersheds in N. Peloponnese. RDI was calculated for several reference periods (3, 6, 9 and 12 months) and SDI was used to represent hydrological drought. Regression equations were derived between RDI and SDI, forecasting the level of hydrological drought for the entire year. Further, DrinC was used for the formulation of a wide range of scenarios representing possible climatic changes and drought events of varying severity in the area of N. Peloponnese. Nomograms were devised for estimating the annual streamflow change caused by each climate change scenario (Tigkas et al. 2012).
- Romania: The Hydrotechnical Engineering Department of the Polytechnics University of Timisoara performed a drought analysis for the western part of Romania, using SPI and RDI, computed with DrinC for several reference periods (Rares 2013).

5. CONCLUDING REMARKS

In this paper the theoretical background, the design and the main features of the Drought Indices Calculator (DrinC) software are presented. The software facilitates the calculation of two recently developed (RDI, SDI) and two widely used drought indices (SPI, PD) and may assist researchers, institutions and stakeholders to derive reliable results for drought analysis.

DrinC can be used in a variety of applications, including drought characterisation, drought monitoring, delineation of the spatial extent of drought, investigation of climatic trends and evaluation of the drought severity for a number of possible drought or future climatic scenarios.

From the applications of DrinC in various regions of the world, it seems that it gradually gains ground for becoming a popular software for drought analysis. Especially, DrinC can be a useful tool in the framework of operational management, assisting for the early estimation of drought consequences and the mitigation of the anticipated impacts.

6. AVAILABILITY AND REQUIREMENTS

The DrinC software can be downloaded from the website <http://drinc.ewra.net>. All the files required for the installation and use of the software (setup file, getting started guide, data input templates etc.) are freely distributed through this site. The software runs on Windows operating systems and it has been successfully tested in all versions (from 98SE to 8.1).

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